



## Observations of molecular oxygen ions in Saturn's inner magnetosphere

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[1] We present an analysis of molecular oxygen ion ( $O_2^+$ ) abundance in Saturn's inner magnetosphere based on observations made with the Cassini Plasma Spectrometer (CAPS). Using data summed over 23 orbits, we resolve and isolate  $O_2^+$  counts from the tail of the water group ion mass distribution ( $W^+$  [ $O^+$ ,  $OH^+$ ,  $H_2O^+$ ,  $H_3O^+$ ]) and from background sources.  $O_2^+$  was initially estimated at  $\sim 1\text{--}2\%$  of the total ion population in the inner magnetosphere based on CAPS data from the Saturn orbital insertion pass. Through refined analysis techniques, we have found  $O_2^+$  to account for just above 0.3% of all  $W^+$  ions at  $L = 4.5$ , with only minor fluctuations in relative density out to  $L = 7.5 - 8.0$ . Beyond  $L = 7.5$ , the relative  $O_2^+/W^+$  abundance exhibits a statistically significant increase through  $L = 10.5$ , which may indicate a neutral  $O_2$  source in the vicinity of Rhea. **Citation:** Martens, H. R., D. B. Reisenfeld, J. D. Williams, R. E. Johnson, and H. T. Smith (2008), Observations of molecular oxygen ions in Saturn's inner magnetosphere, *Geophys. Res. Lett.*, 35, L20103, doi:10.1029/2008GL035433.

### 1. Introduction

[2] The Cassini spacecraft has been orbiting Saturn for over four years, returning far more data than the Pioneer and twin Voyager spacecraft combined, in an unprecedented mission to study the planet, its satellites, and its magnetosphere. Cassini's long-term, direct observation of Saturn's plasma environment has allowed a rigorous investigation into magnetospheric composition, including the presence and source of minor ion populations (i.e.,  $<5\%$  of the total ion population). Analysis of these minor species can provide clues to the origin and evolution of Saturn's rings and satellites. Here we focus our investigation on molecular oxygen ions ( $O_2^+$ ) in Saturn's inner magnetosphere ( $L\text{-shell} \leq 11$ ).

[3] Prior to Cassini, laboratory experiments and Hubble Space Telescope observations suggested the presence of water products in Saturn's magnetosphere [e.g., Johnson *et al.*, 1989; Shemansky *et al.*, 1993; Jurac *et al.*, 2002], and ozone ( $O_3$ ) in the icy surfaces of Dione and Rhea [Noll *et al.*, 1997]. During Saturn Orbital Insertion (SOI), Cassini detected an ionosphere dominated by  $O_2^+$  and  $O^+$  over Saturn's main rings [Tokar *et al.*, 2005; Young *et al.*,

2005; Waite *et al.*, 2005]. Johnson *et al.* [2006] determined that these ions are photo-ionization products of neutral  $O_2$ , which is produced primarily through photo-induced decomposition of ring particles into  $H_2$  and  $O_2$ . Radiation-induced decomposition can also occur via high-energy charged-particle impacts on ice grains in the tenuous rings and icy satellite surfaces [Johnson *et al.*, 2006]. Due to the relatively small gravity of Saturn's satellites and rings, the decomposed ice particles form toroidal oxygen atmospheres as observed above the main rings [Johnson *et al.*, 2006; Tokar *et al.*, 2005] and as predicted at the icy satellites [e.g., Sittler *et al.*, 2004] and tenuous E, F, and G rings.

[4] Using Cassini Ion Mass Spectrometer (IMS) data collected during SOI, Sittler *et al.* [2005] estimated the  $O_2^+$  abundance in Saturn's inner magnetosphere ( $4 \leq L \leq 14$ ) at roughly  $1\text{--}2\%$  of the total ion population. Johnson *et al.* [2005] described how the  $O_2$  from the rings may attain relatively large energies and be redistributed throughout the magnetosphere before becoming ionized. The question of whether detections of  $O_2^+$  outside the main rings indicate additional local sources of  $O_2$  remains open.

[5] In this paper we present an analysis of  $O_2^+$  abundances in Saturn's inner magnetosphere using Cassini Plasma Spectrometer (CAPS) IMS data collected over 23 orbits of the Cassini spacecraft. We consider whether the  $O_2^+$  distribution suggests additional sources of  $O_2$ . Due to the relative weakness of the  $O_2^+$  signal, special attention must be paid to the method by which quantitative information is extracted. Thus, we first describe in detail the analysis technique, followed by a discussion of our results.

### 2. Data Collection

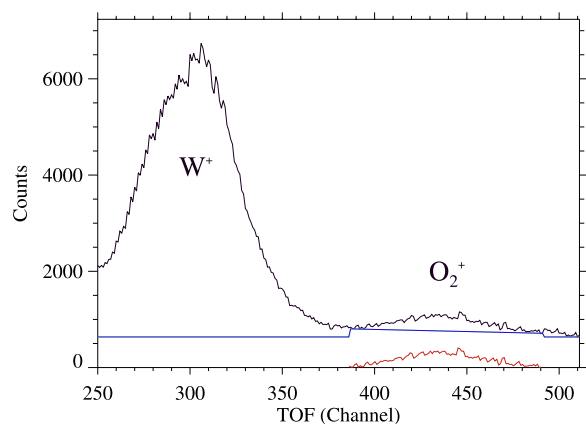
[6] The CAPS IMS consists of a top-hat style electrostatic analyzer, followed by a linear electric field time-of-flight (TOF) mass analyzer. TOF spectra are collected as a function of analyzer voltage, yielding two-dimensional energy-versus-TOF arrays. (See Young *et al.* [2004] for details on IMS operation.) Mass analysis is facilitated by two detector systems: the high-resolution ( $M/\Delta M \sim 60$ ) linear electric field (LEF) detector and the lower resolution ( $M/\Delta M \sim 8$ ) straight-through (ST) detector. Despite lower mass resolution, the ST detector boasts higher detection efficiencies than the LEF detector, which is useful for amplifying the weak signals from minor magnetospheric constituents. Hence, data obtained by the ST detector are used exclusively in our investigation.

[7] TOF spectra for 23 orbits between the dates of 27 October 2004 and 27 March 2006 were first summed and then binned by distance from Saturn for every 0.5 dipole L-shell interval from  $L = 3.5$  to  $L = 14$ . Periapses of

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**Figure 1.** Time-of-flight spectrogram for  $L = 5.5 - 6.0$  at energy 666.4 eV, zoomed in to focus on the  $O_2^+$  peak. The black line indicates the original counts prior to background subtraction. The blue indicates the background subtraction. The red line depicts the difference between the black and blue lines and represents the final set of counts to be applied to our  $O_2^+$  summation method.

these orbits fall within the dusk-side portion of the magnetosphere, generally between 16 and 20 local time. Although these orbits span  $\pm 22^\circ$  of latitude, over half of them were equatorial.

[8] By summing large amounts of data we can amplify a weak  $O_2^+$  signal over a significant background radiation signal and/or overlap with the spectral tails of the dominant ion population, the water group ions ( $O^+$ ,  $OH^+$ ,  $H_2O^+$ , and  $H_3O^+$ ; or collectively,  $W^+$ ). This allows us to detect and quantify  $O_2^+$  abundances from  $L = 4.5$  to  $L = 11$ . Although we also detect  $O_2^+$  inward to  $L = 3.5$ , penetrating radiation overwhelms the  $O_2^+$  signal for  $L < 4.5$ , significantly reducing our detection confidence for this region and preventing useful analysis. Likewise,  $O_2^+$  is qualitatively observed outward to at least  $L = 12$ ; however, beyond  $L = 11$ , the signal weakens significantly and soon drops below the detection limits of the IMS.

[9] Inward of the plasmasphere boundary at  $L \sim 14$ , we expect the plasma to be nearly co-rotating [Young et al., 2005; Sittler et al., 2006; Wilson et al., 2008], but, as CAPS is not always directed toward the co-rotational plasma flow, data filtration was necessary to minimize the effect of plasma anisotropy on the composition analysis. Hence, we removed data obtained when the direction of co-rotation was  $30^\circ$  beyond the CAPS field of view as well as when CAPS was not actuating. (The CAPS actuator rotates the IMS aperture through  $220^\circ$  to compensate for CAPS being mounted on a non-spinning spacecraft.) Also, to avoid contaminating the dataset with composition conditions localized to a particular satellite's environment, all satellite encounters are explicitly excluded.

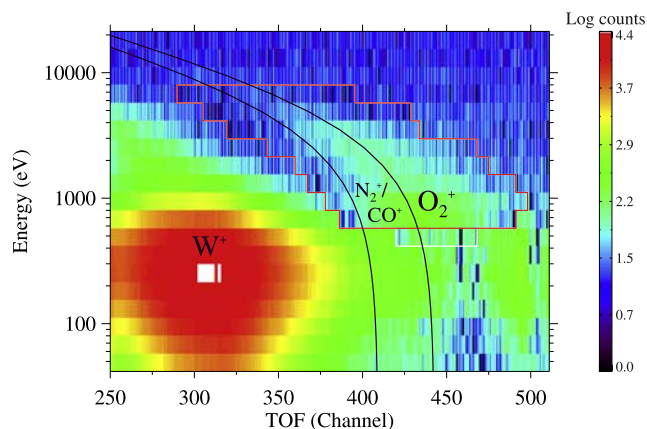
### 3. Species Isolation

[10] To determine  $W^+$  abundance, we used an automated analysis method whereby model peaks derived from IMS calibration were fit to the TOF spectra via a linear least-squares algorithm. Dependable for dominant ion populations, this method has been routinely applied to the analysis

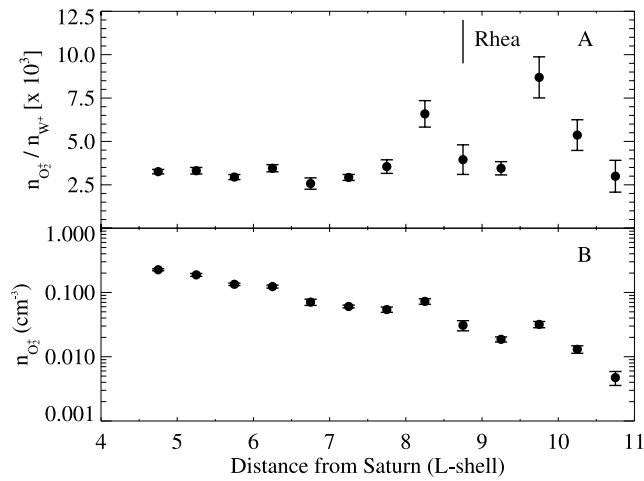
of IMS composition data [Young et al., 2005; Tokar et al., 2005; Sittler et al., 2006]. However, the procedure becomes less reliable when dealing with minor ion populations because, although the model functions accurately fit the central portion of peaks, uncertainties become significant in the far tails due to differences in the calibration and space environments. Hence, when subtracting the model function for a major species (e.g.,  $H_2O^+$ ) to reveal the presence of a minor species (e.g.,  $O_2^+$ ) that appears in the tail, the potential exists to miscount the content of the minor species, requiring the use of an alternative isolation method.

[11] Hence, for a given L-shell bin of our summed TOF spectra,  $O_2^+$  was isolated using a multi-step process to remove background and  $W^+$  tail counts. First, a flat background correction eliminated the majority of background counts. Next, TOF spectra for certain energies, typically between  $\sim 400$ – $2000$  eV, required the subtraction of a sloped baseline to account for the  $W^+$  tail, which could not be accurately determined by the automated fitting method previously described. In such cases, the boundaries of the  $O_2^+$  peak were determined by empirically fitting the mass peaks and finding intersection points between the  $W^+$  and  $O_2^+$  peak tails. Next, the minima within 10 channels on either side of both  $O_2^+$  boundary points were determined; one standard deviation of the count values was added to avoid underestimating the  $W^+$  tail contamination (i.e., to determine the average, not minimum, background); and the resultant line connecting the two points was subtracted from the isolated  $O_2^+$  peak. Figure 1 depicts this process graphically for the summed TOF spectra between  $L = 5.5 - 6.0$ .

[12] Figure 2 is an example of the summed TOF spectrogram for  $L = 5.5 - 6.0$ , restricted to a TOF range of 250–511 channels, and an energy range of 41.6–21,300 eV. The spectrogram, plotted after the removal of background counts, illustrates the prominent  $W^+$  population as well as the region selected for  $O_2^+$  count integration, depicted by the red box. Despite the relative weakness of the  $O_2^+$  signal, it is well resolved from  $W^+$  because  $O_2^+$  peaks at energies a



**Figure 2.** Time-of-flight spectrogram depicting the summation of counts over the analysis period for  $L = 5.5 - 6.0$ . The curved, black lines indicate the calculated peak centers for (left)  $N_2^+/CO^+$  and (right)  $O_2^+$ . The red box indicates the determined boundaries for the  $O_2^+$  peak. The white box extension represents a more uncertain  $O_2^+$  detection, included as a source of error.



**Figure 3.**  $O_2^+$  number densities: (a) Density ratio of  $O_2^+$  to  $W^+$  as a function of L-shell distance from Saturn. Error is quoted at the  $1-\sigma$  level. (b) Absolute  $O_2^+$  densities determined by scaling our  $O_2^+/W^+$  density ratio by *Wilson et al.*'s [2008] parameterized fit to  $W^+$ .

factor of 2–3 higher than  $W^+$ . However, the  $O_2^+$  count contribution does become difficult to resolve at lower energies where the  $W^+$  counts peak (white-bounded extension in Figure 2). The uncertainty from isolating counts at these energies dominates the error estimate for  $O_2^+$  abundance. Additional sources of uncertainty arise from counting statistics and uncertainty in the detection efficiency. Note that because each radial bin is composed of counts summed over multiple orbits, the error bars reflect an uncertainty in the mean and not variability in the environment.

[13] *Smith et al.* [2008] have reported the presence of a small mass-28 population in Saturn's inner magnetosphere as well (i.e.,  $N_2^+$  and/or  $CO^+$ ), contained within the left shoulder of the  $O_2^+$  peak. The mass-28 population is difficult to quantify, however, due to limited mass resolution and its proximity to  $O_2^+$ . Fitting calibration-determined model functions to the data reveals an average  $(N_2^+ + CO^+)/O_2^+$  relative abundance of  $\sim 2\%$  over the investigation region (consistent with the rough upper limits placed on  $N_2^+/CO^+$  by *Smith et al.* [2008]). However, because of the large uncertainty in the fitting procedure and the small size of the contribution, we have chosen not to correct the  $O_2^+$  abundance for the mass-28 contribution, but do include it as a source of uncertainty.

#### 4. Results and Discussion

[14] Figure 3a presents our calculated  $O_2^+/W^+$  relative number densities as a function of distance from Saturn. The mass of the collective  $W^+$  distribution was approximated at 17 amu. Figure 3b depicts an estimate of absolute  $O_2^+$  number density, calculated by scaling the  $O_2^+/W^+$  ratio by a parameterized fit to the equatorial water group density ( $n_{W^+} = 162.1 \exp(-0.042L^2)$ ), as reported by *Wilson et al.* [2008].

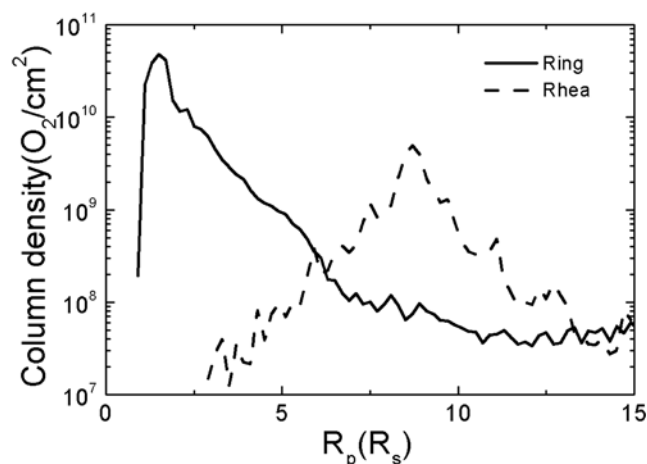
[15] Our analysis indicates that, on average,  $O_2^+$  number density is just above 0.3% relative to  $W^+$  at  $L = 4.5$ , with only minor fluctuations out to  $L = 7.5 - 8.0$ . Since  $O_2^+$  and  $W^+$  are close in mass, the relatively flat density ratio for  $L = 4.5 - 7.5$  is consistent with sources for both species

being primarily interior to  $L = 4.5$  (i.e., the rings for  $O_2^+$  [*Tokar et al.*, 2005] and Enceladus for  $W^+$  [*Tokar et al.*, 2006]). Beyond  $L = 7.5$ , however, the relative density ratio increases. For  $L = 4.5 - 7.5$ , the average relative density ratio is  $0.31 \pm 0.01\%$ ; for  $L = 7.5 - 10.5$ , it is  $0.41 \pm 0.02\%$ . Thus, the increase in  $O_2^+$  abundance beyond  $L = 7.5$  is statistically significant, suggesting the possibility of additional  $O_2^+$  sources or a change in the ionization and/or loss rate. Additional sources might include local E-ring grains and, since Rhea orbits at  $\sim 8.75R_S$ , the surface of Rhea and its newly inferred ring particles [*Jones et al.*, 2008].

[16] The energetic protons and electrons present in Saturn's magnetosphere are relatively inefficient sputtering agents, but cause the decomposition of icy surfaces into  $O_2$  and  $H_2$  [*Johnson*, 1996; *Johnson and Quickenden*, 1997], contributing to the neutral population in Saturn's magnetosphere [*Sittler et al.*, 2004]. Since the density of ice grains in Saturn's E-ring increases inward of Rhea's orbit [*Showalter et al.*, 1991], and the density of energetic protons decreases rapidly inside of  $\sim 8R_S$  [*Paranicas et al.*, 2008], an enhancement in the region seen in Figure 3a is possible due to decomposition of E-ring grains. However, the total surface area of the E-ring at this orbital region is orders of magnitude smaller than the surface area of Rhea, Saturn's largest icy satellite. In addition, the energetic electron population ( $\sim 10$ 's of keV) exhibits strong depletions at Rhea's orbit and over a region on the order of a few Rhea radii [*Jones et al.*, 2008], suggesting that energetic charged particles are depositing energy into Rhea's surface and its ring. Presuming Rhea's surrounding material is primarily icy, like Rhea's surface, the energy deposited will be a source of  $O_2$ . The decomposition of ice by particle radiation is corroborated by earlier observations of  $O_3$  trapped in Rhea's surface ice [*Noll et al.*, 1997]. Therefore, we suggest this additional source of  $O_2$  produces our observed enhancement in relative  $O_2^+$  abundance.

[17] We estimated a rough source rate for  $O_2$  production from Rhea and its putative trapped ice grains using the observed depletions in energetic electron flux [*Jones et al.*, 2008]. Due to the considerable uncertainty in  $O_2$  production efficiency, the source rate should be considered an order-of-magnitude estimate. Ejected molecules are tracked in Saturn's gravitational field until they are lost by ionization or charge exchange [e.g., *Johnson et al.*, 2004; *Sittler et al.*, 2008]. Figure 4 compares the local  $O_2$  column density to the column of  $O_2$  that has been scattered out of Saturn's ring atmosphere [*Johnson et al.*, 2006]. Although both are order-of-magnitude estimates, an enhancement in  $O_2$  density, implying a relative increase in  $O_2^+$ , is consistent with a Rhea source. In addition, the spatial extent of the enhancement in the  $O_2$  column density is consistent with the ion data. The simulated neutral column densities, however, are well below the upper limits established by MIMI ENA and UVIS observations [*Jones et al.*, 2008], and therefore could only be detected as a source of plasma ions.

[18] Although this enhancement is roughly consistent with our observations, it does not show the apparent drop in  $O_2^+$  relative abundance at Rhea's orbit, as seen in Figure 3a. Since the enhancement in  $O_2^+$  is locally produced, whereas  $W^+$  is primarily a product of outward diffusion at this orbital radius [*Sittler et al.*, 2008], the drop may be due to the large target presented by Rhea's gravitationally bound



**Figure 4.** Simulations of the  $O_2$  neutral column density as a function of radial distance from Saturn (1 Saturn radii [ $R_s$ ] = 60330 km). Solid line:  $O_2$  scattered from over the main rings by ion-molecule collisions [Johnson et al., 2006]: source rate (S)  $\sim 4 \times 10^{26}$   $O_2/s$ ; loss is through re-impacting main rings; ejecta speed distribution is due to ion-neutral scattering; over the main rings the total column of  $O_2$  is the scattered component shown here plus the unscattered component ( $\sim 10^{13}$   $O_2/cm^2/s$  [Johnson et al., 2006]). Dashed line:  $O_2$  from Rhea and its ring material, assumed to be water ice. S is estimated from energy deposited by energetic particles [Jones et al., 2008] and lifetime,  $\tau$ , is taken from Sittler et al. [2008]:  $S\tau \sim 2 \times 10^{31}$  particles. These sources are order-of-magnitude estimates. Although the source rate from the vicinity of Rhea is 2–3 orders of magnitude smaller, the ejected  $O_2$  survive about 3 orders of magnitude longer since those ejected from the rings have a very high probability of re-impact in one orbit.

grains to the fresh pick-up ions. That is, the gravitationally bound grains preferentially absorb  $O_2^+$  ions, which have very large gyro radii, at a faster rate than the outward diffusing  $W^+$  ions.

## 5. Summary and Conclusions

[19] Using CAPS IMS data, we have analyzed molecular oxygen ion abundance in Saturn's inner magnetosphere. On average,  $O_2^+$  number density is just above 0.3% relative to  $W^+$  at  $L = 4.5$ , with only minor fluctuations out to  $L = 7.5$ . This marks the transition to a statistically significant increase in relative density out to  $L = 11$ , which may indicate that Rhea and its putative ring are sources of neutral  $O_2$ , and hence  $O_2^+$ , in the inner magnetosphere. Although the observed enhancement here could be consistent with a local increase in the ionization rate or a radial variation in the proton flux, we suggest that the enhancement is consistent with recent observations that indicate large depletions in energetic electrons within Rhea's Hill sphere. Such energetic particles are a source of  $O_2$  by radiation-induced decomposition of icy materials. It also appears that  $O_2^+$  produced throughout the toroidal cloud are being absorbed preferentially by Rhea and its cloud of ice grains, accounting for the apparent drop in  $O_2^+/W^+$  relative density between  $L = 8.5$  and  $L = 9.5$ .

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